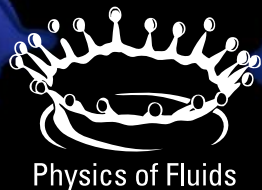


# Convection-driven porous media flows: Implications for carbon dioxide sequestration



M. De Paoli<sup>1,2</sup>

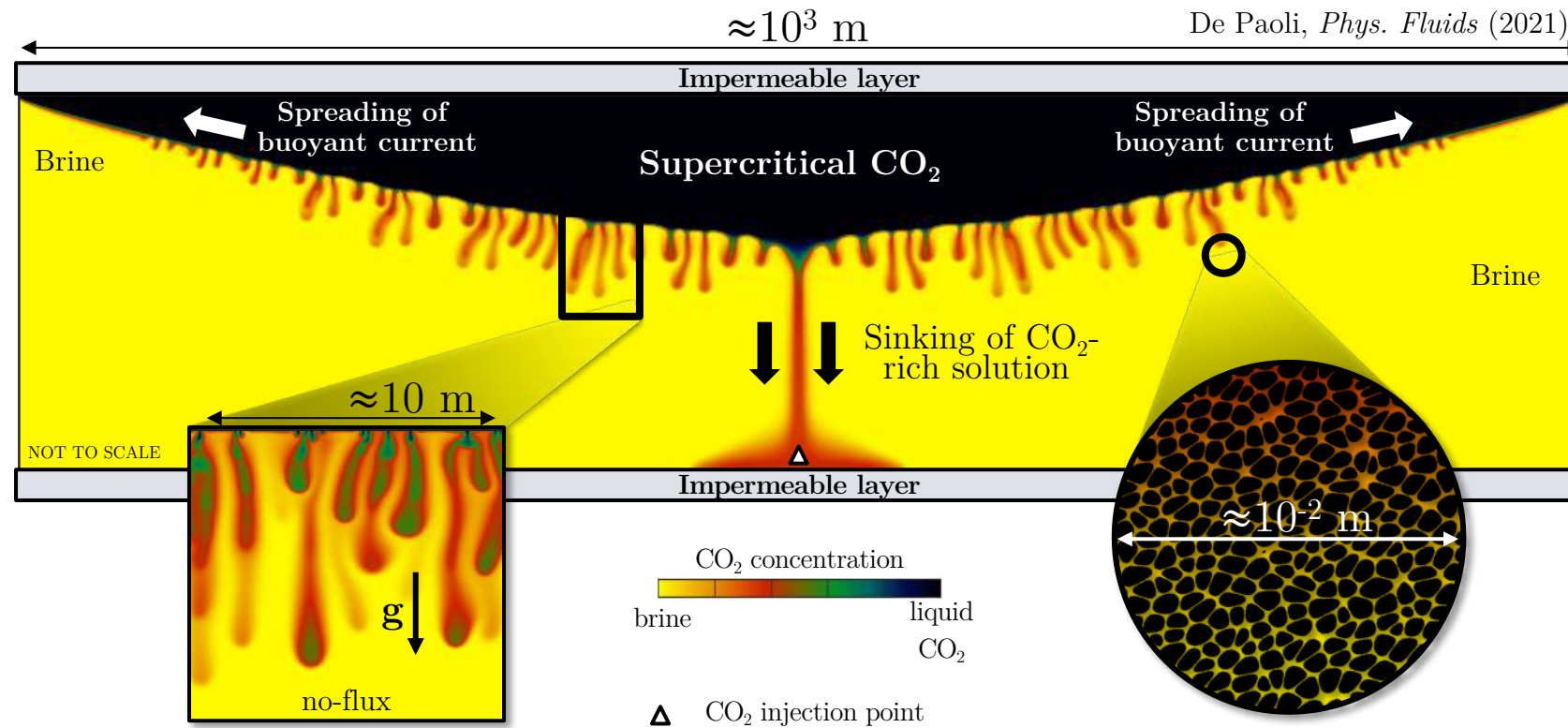
[m.depaoli@utwente.nl](mailto:m.depaoli@utwente.nl)

<sup>1</sup>Physics of Fluids Group, University of Twente, Enschede (The Netherlands)

<sup>2</sup>Institute of Fluid Mechanics and Heat Transfer, TU Wien, Vienna (Austria)



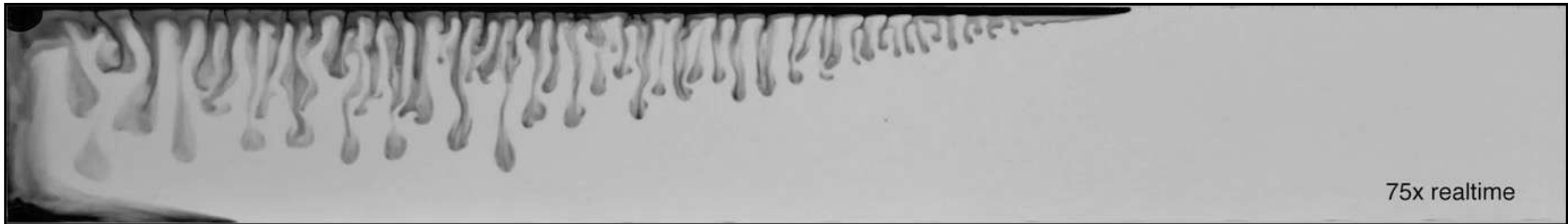




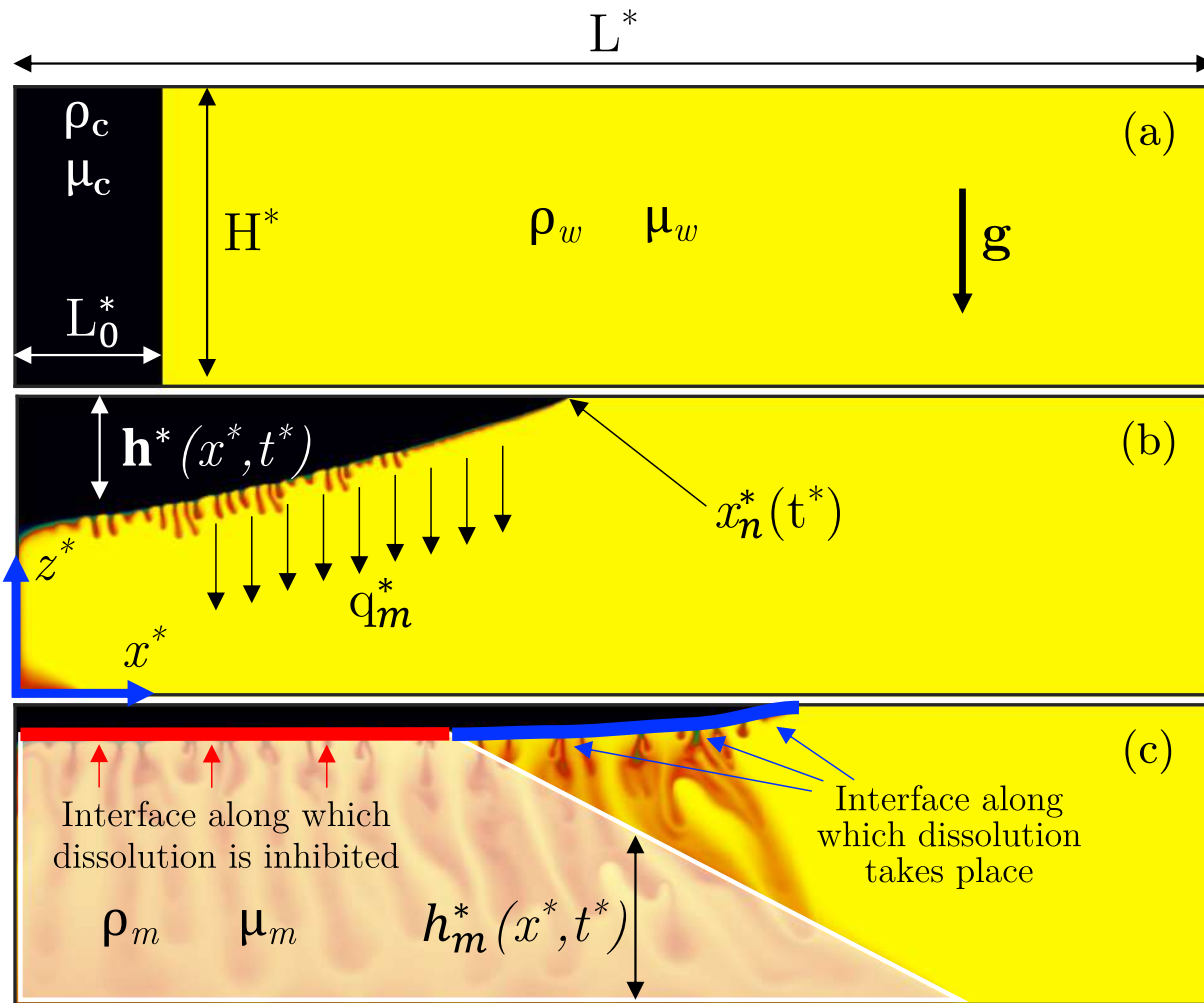
## Reservoir properties

- anisotropy and heterogeneities
- finite size of confining layers
- effects of rock properties (mechanical dispersion)
- chemical dissolution and morphology variations
- ...

MacMinn et al., *Geophys. Res. Lett.* (2013)



# Multiphase gravity currents with dissolution



De Paoli, *Phys. Fluids*. (2021)

$$\frac{\partial h}{\partial t} - \frac{\partial}{\partial x} \left[ (1-f)h \frac{\partial h}{\partial x} - \delta f h_m \frac{\partial h_m}{\partial x} \right] = -\varepsilon_0,$$

$$\frac{\partial h_m}{\partial t} - \frac{\partial}{\partial x} \left[ \delta(1-f_m)h_m \frac{\partial h_m}{\partial x} - f_m h \frac{\partial h}{\partial x} \right] = \frac{\varepsilon_0}{X_v}$$

$$f = \frac{Mh^*/H^*}{(M-1)h^*/H^* + (M_m-1)h_m^*/H^* + 1},$$

$$f_m = \frac{M_m h_m^*/H^*}{(M-1)h^*/H^* + (M_m-1)h_m^*/H^* + 1},$$

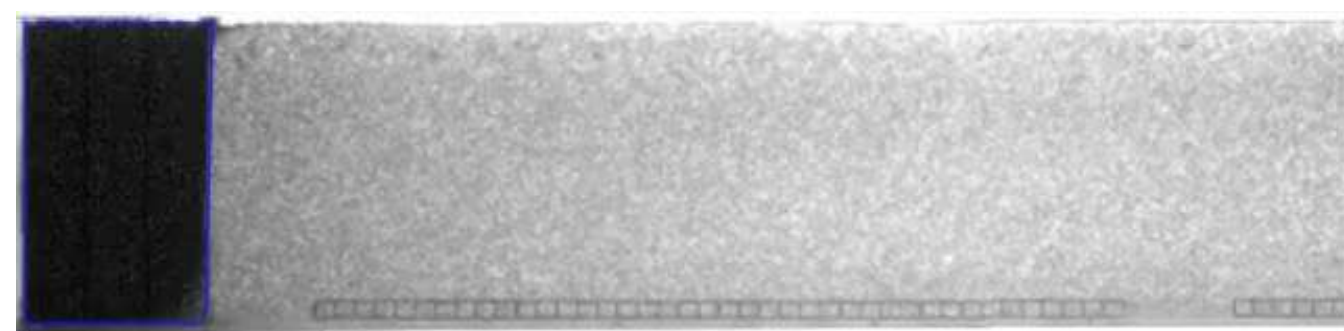
MacMinn, Neufeld, Hesse,  
and Huppert, *Water Resour. Res.* (2012)

Mobility ratios  $M = \mu_w/\mu_c$  and  $M_m = \mu_w/\mu_m$

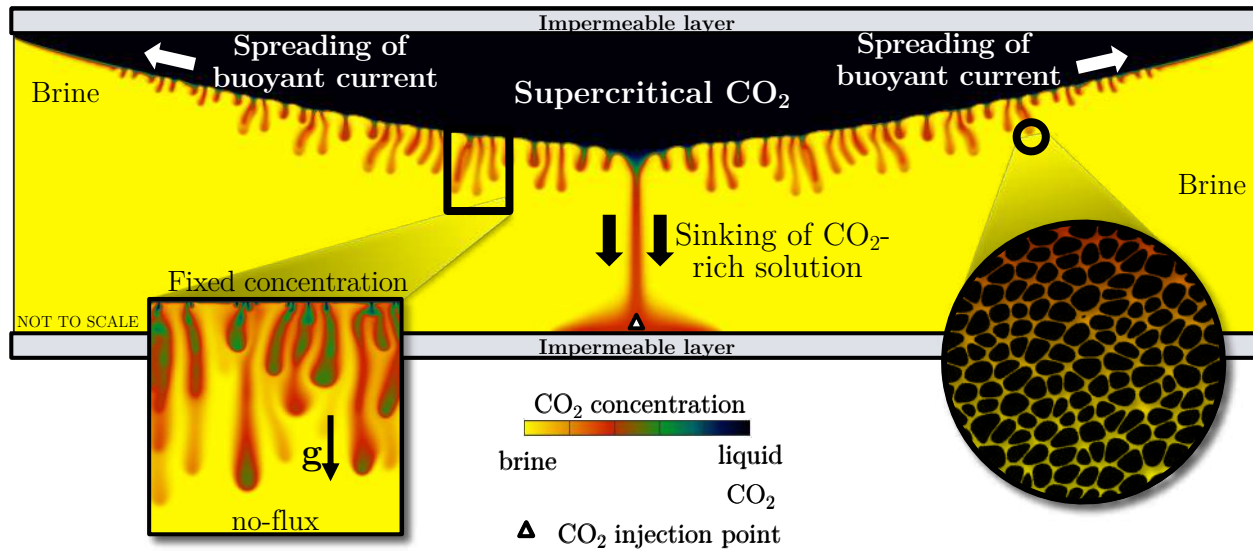
Buoyancy velocity ratio  $\delta = W_m^*/W^*$

Volume fraction  $X_v = \rho_m X_m / \rho_c$

# Multiphase gravity currents with dissolution



MacMinn, Neufeld, Hesse, and Huppert, *Water Resour. Res.* (2012)



$$\frac{\partial h}{\partial t} - \frac{\partial}{\partial x} \left[ (1-f)h \frac{\partial h}{\partial x} - \delta f h_m \frac{\partial h_m}{\partial x} \right] = -\varepsilon_0$$

$$\frac{\partial h_m}{\partial t} - \frac{\partial}{\partial x} \left[ \delta(1-f_m)h_m \frac{\partial h_m}{\partial x} - f_m h \frac{\partial h}{\partial x} \right] = \frac{\varepsilon_0}{X_v}$$

$$f = \frac{Mh^*/H^*}{(M-1)h^*/H^* + (M_m-1)h_m^*/H^* + 1},$$

$$f_m = \frac{M_m h_m^*/H^*}{(M-1)h^*/H^* + (M_m-1)h_m^*/H^* + 1},$$

$$\varepsilon_0(x) = \begin{cases} 0 & \text{if } h(x) = 0 \text{ or } h(x) + h_m(x) = 1 \\ \varepsilon & \text{else,} \end{cases}$$

$$\varepsilon = \frac{q_m^*}{\phi W^*} \left( \frac{L_0^*}{H^*} \right)^2$$

How to determine the dissolution rate  $q_m^*$  ?

## Dimensionless equations

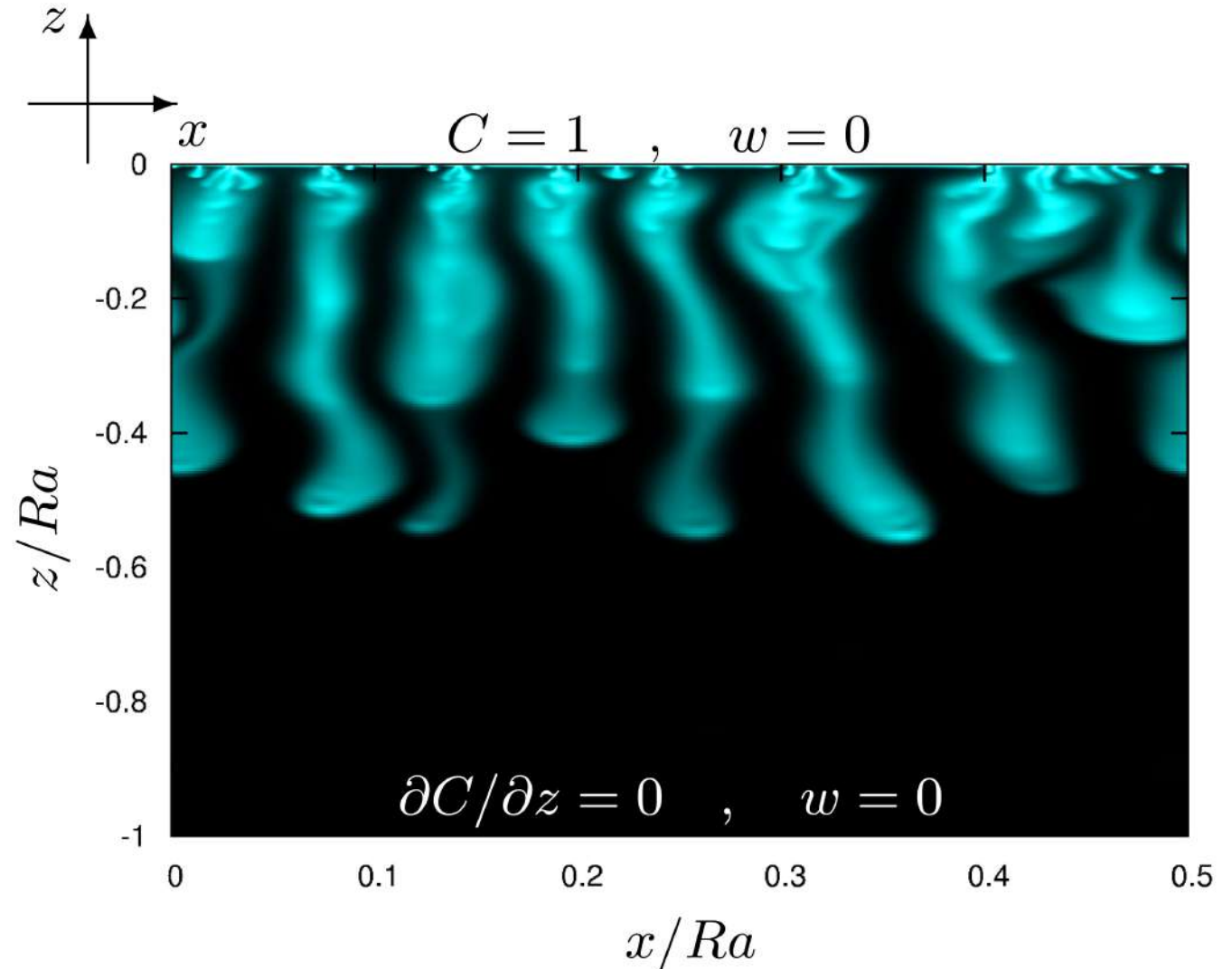
$$\frac{\partial C}{\partial t} + u \frac{\partial C}{\partial x} + w \frac{\partial C}{\partial z} = \frac{1}{Ra} \left( \frac{\partial^2 C}{\partial x^2} + \frac{\partial^2 C}{\partial z^2} \right)$$

$$u = -\frac{\partial P}{\partial x} \quad , \quad w = -\frac{\partial P}{\partial z} - C$$

$$\frac{\partial u}{\partial x} + \frac{\partial w}{\partial z} = 0$$

## Governing parameter

$$Ra = \frac{gH^*k_v\Delta\rho^*}{\mu\Phi D}$$



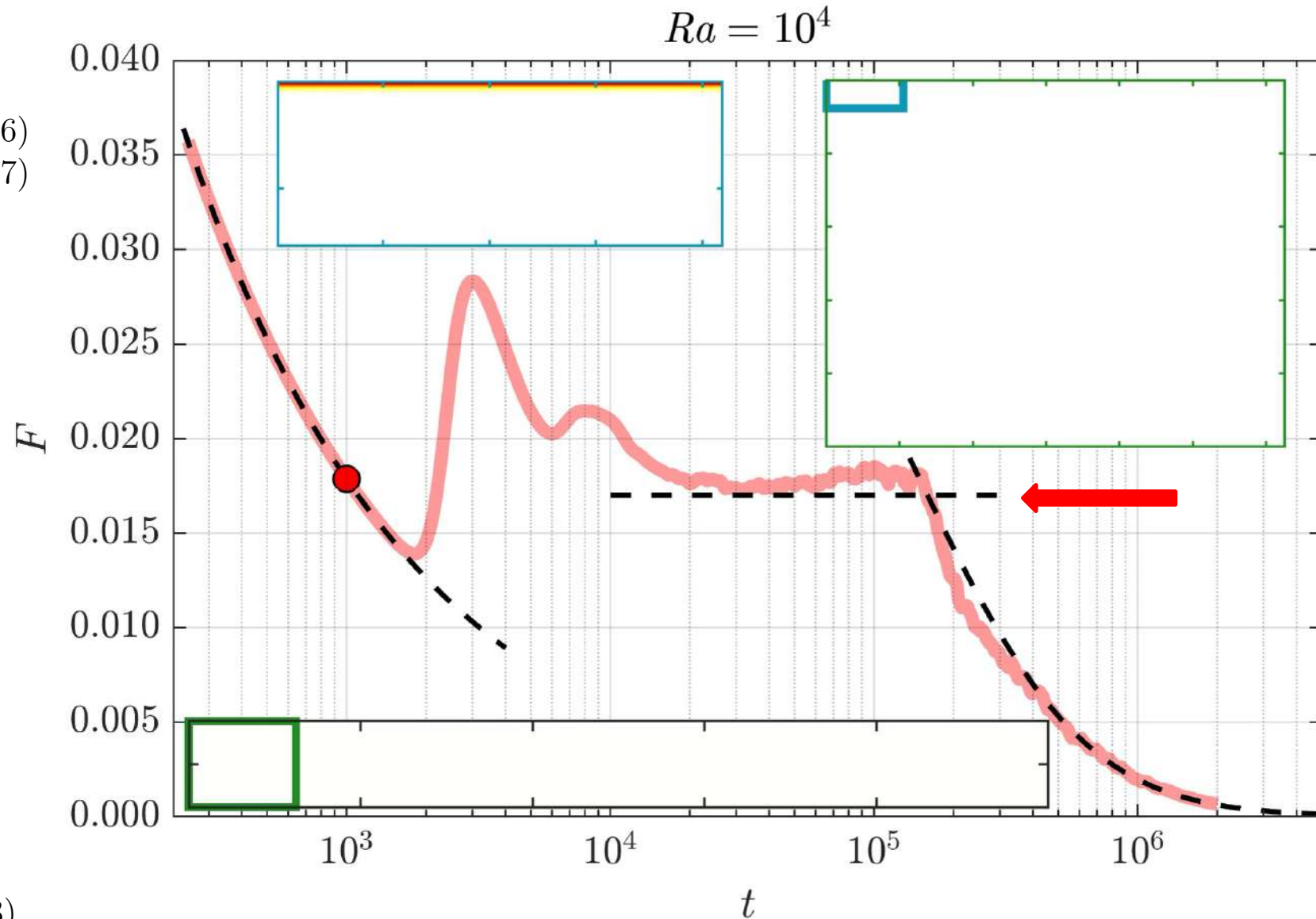


De Paoli, Zonta and Soldati, *Phys. Fluids* (2016)  
De Paoli, Zonta and Soldati, *Phys. Fluids* (2017)

$$F(t) = \frac{1}{L} \int_0^L \left. \frac{\partial C}{\partial z} \right|_{z=0} dx$$

Examples of model extension:  
effect of **anisotropy** of the medium

See also Slim, *J. Fluid Mech.* (2014)  
Hewitt, Neufeld & Lister, *J. Fluid Mech.* (2013)



# Convection in anisotropic media

Examples of model extension:  
effect of **anisotropy** of the medium

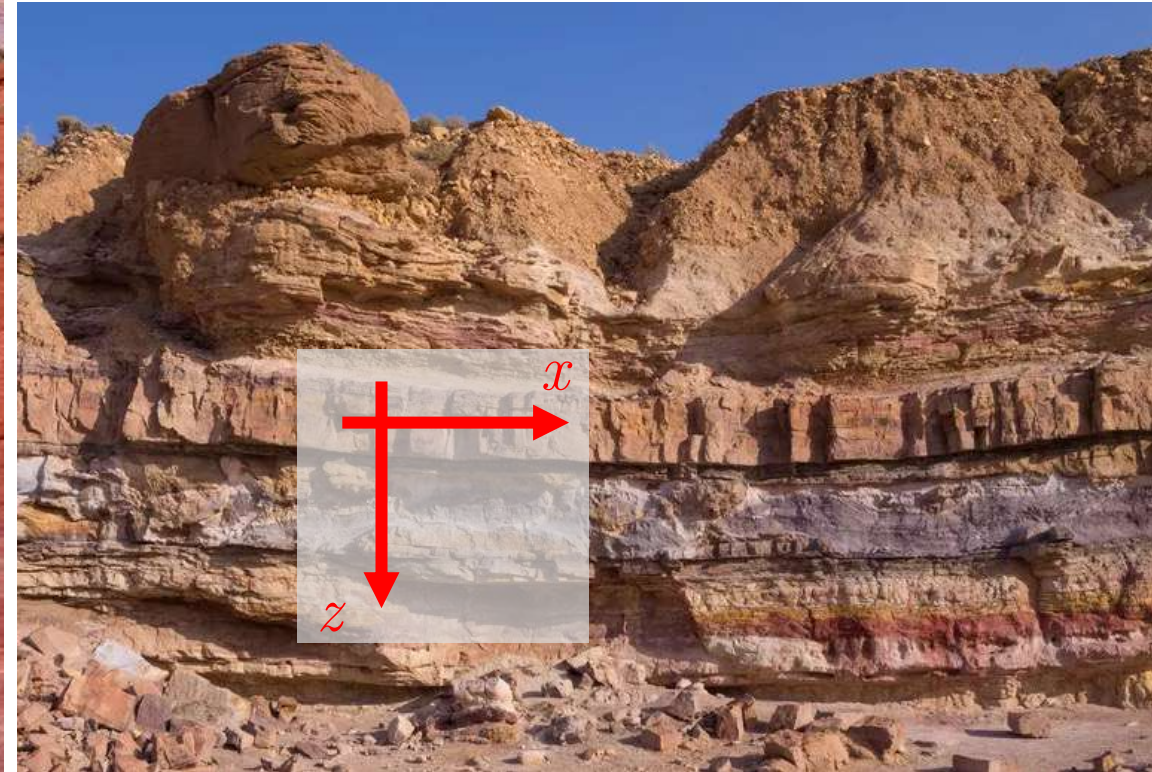
Sedimentary rocks: Rocks formed by stratification

Assumptions:

1. Homogeneous porous medium
2. Anisotropic porous medium
  - Principal directions of the permeability tensor aligned with the reference frame



benedek / Getty Images

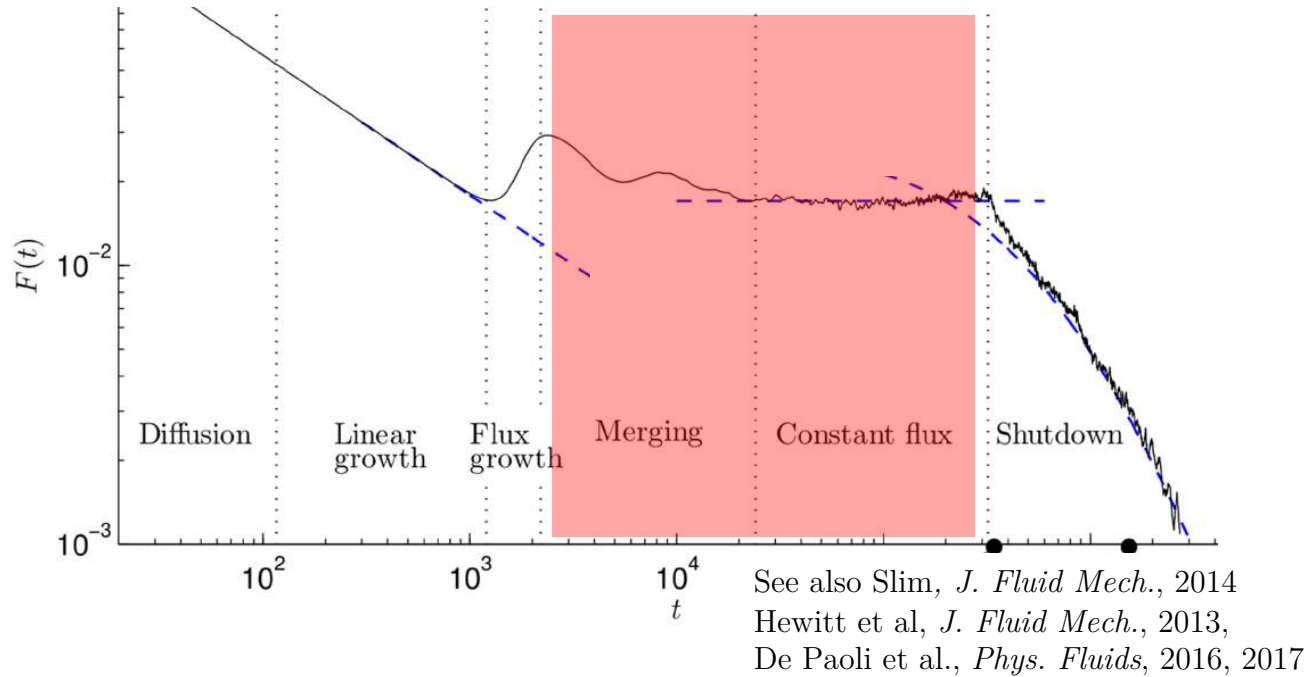


Rhododendrites/Wikimedia Commons/CC BY 4.0



# Heat flux and dissolution rate

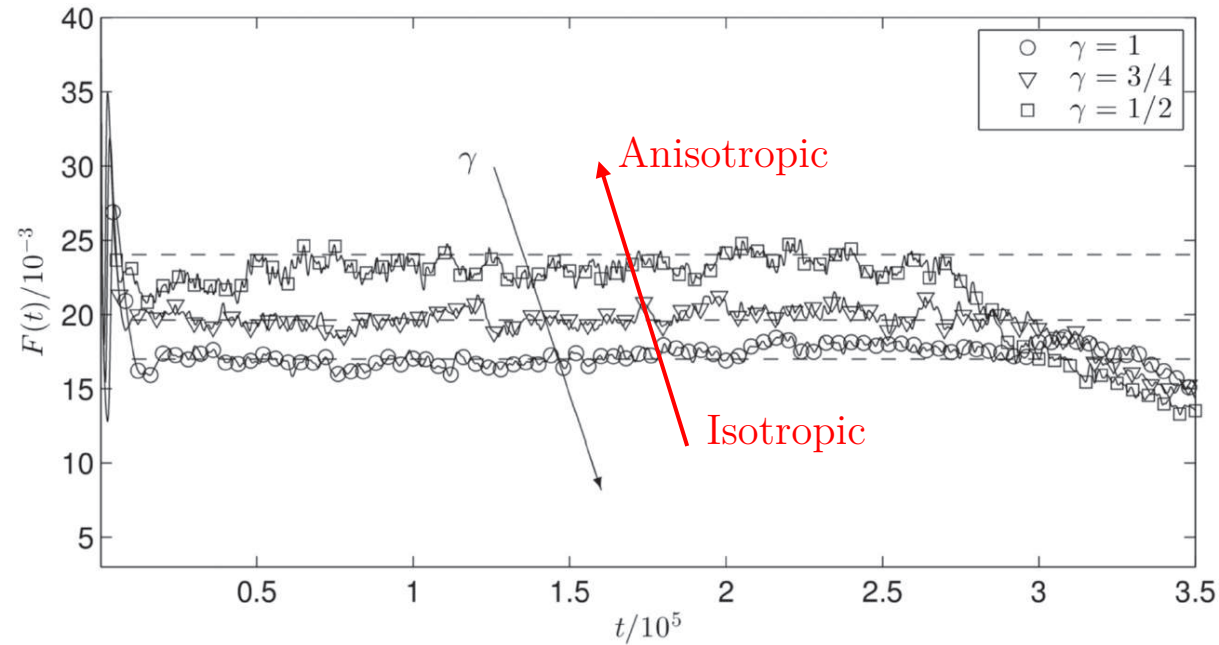
isotropic medium



Convection-dominated  $F(t) = 0.017$

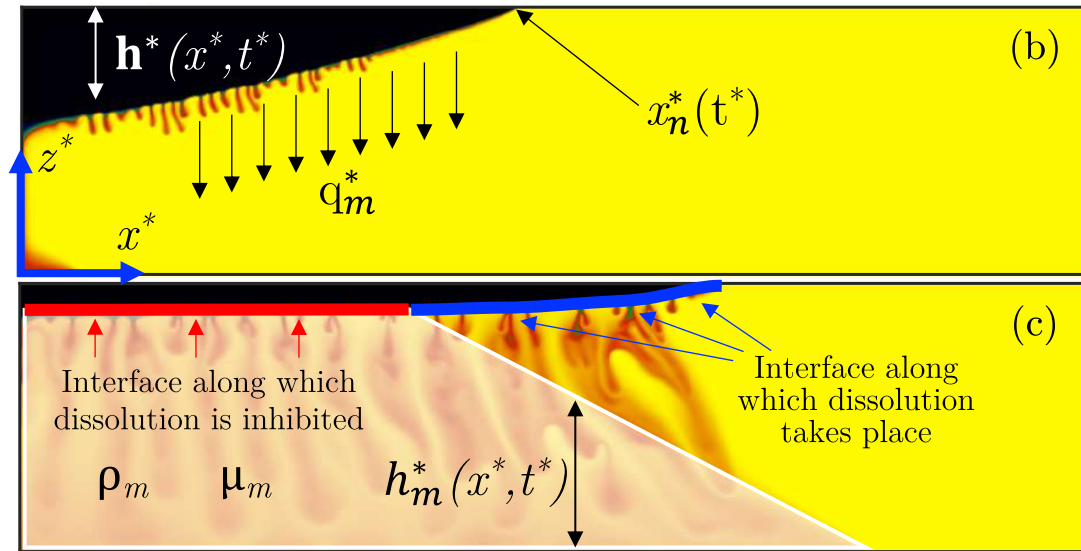
Strong influence of  $\gamma$  on flux

$$F(t) = \frac{1}{L} \int_0^L \frac{\partial C}{\partial z} \Big|_{z=0} dx$$



$$q_m^* \equiv F(t) = 0.017\gamma^{-1/2}$$





Darcy-scale simulations:



$$\text{dissolution rate } q_m^* \sim \gamma^{-\frac{1}{2}}$$



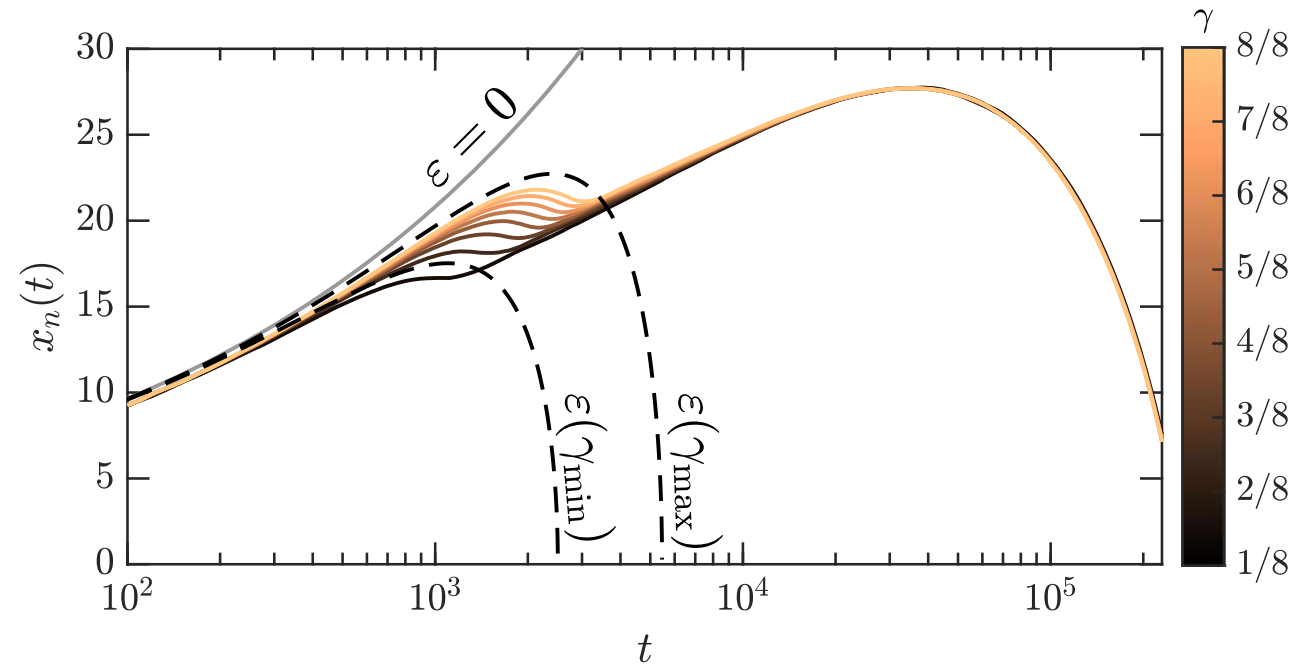
dissolution increases with the anisotropy of the medium

Sedimentary rocks are anisotropic

$$\gamma = \frac{k_v}{k_h} < 1$$

$\gamma = 1$  isotropic

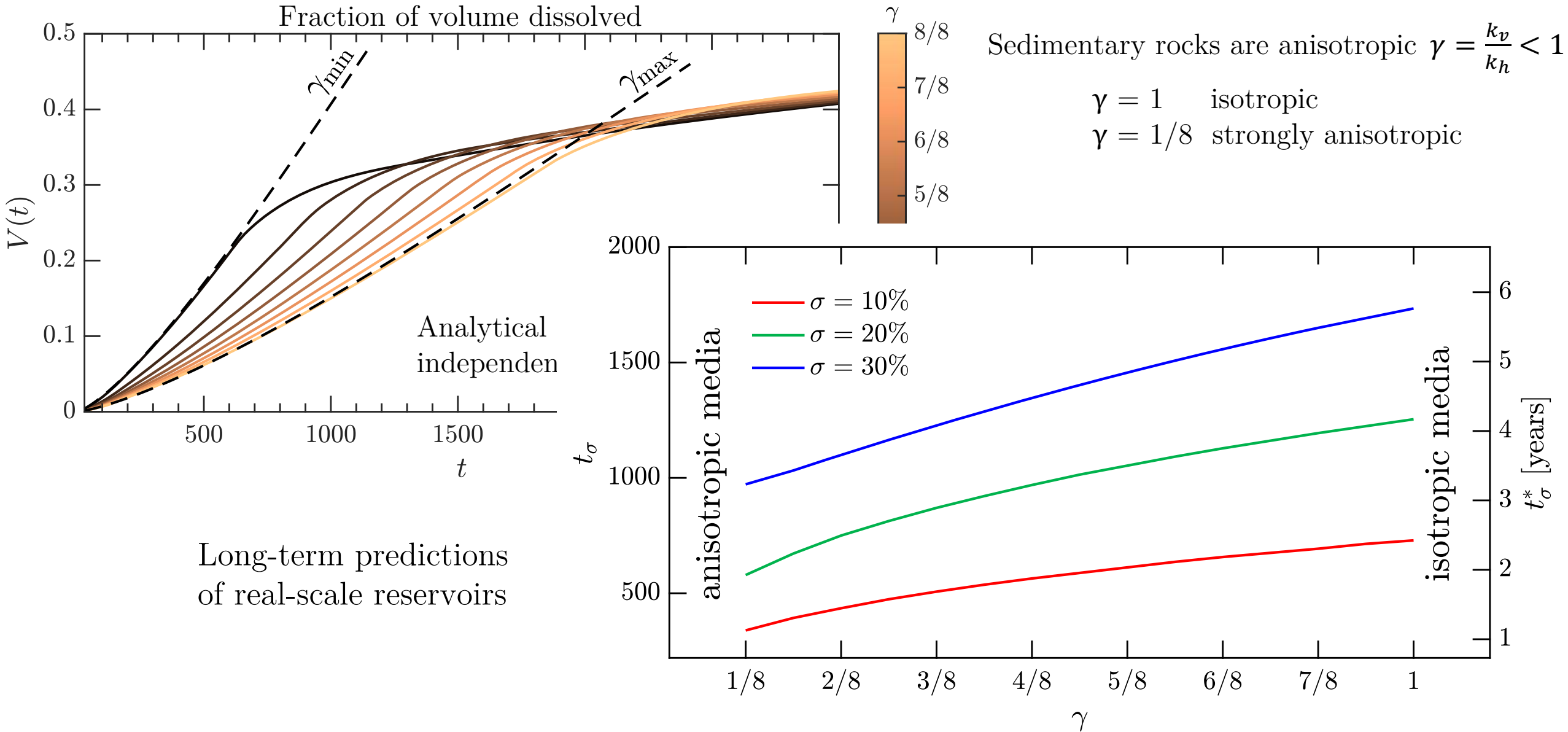
$\gamma = 1/8$  strongly anisotropic



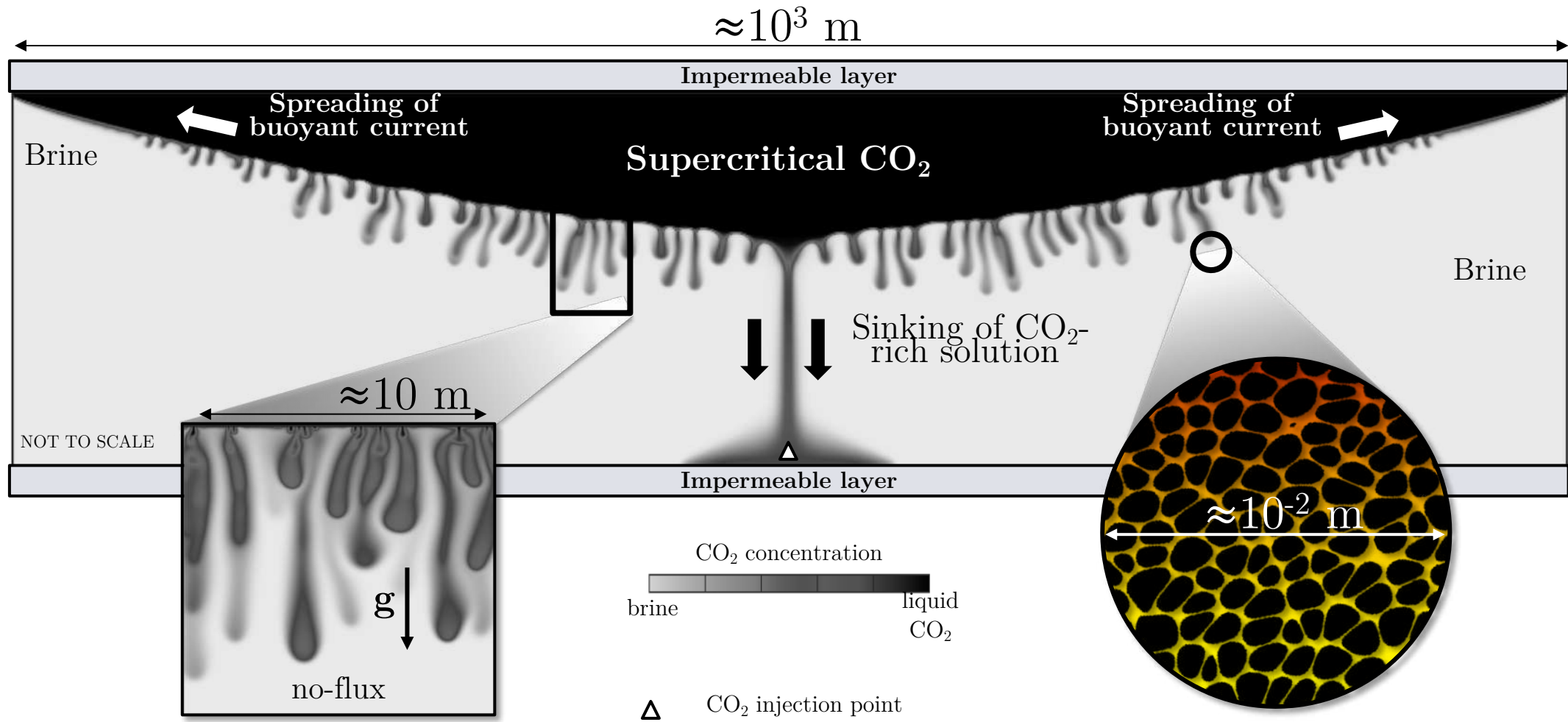
Analytical solution in case of

- no-dissolution —————
- independent currents - - - - -

# Effect of anisotropy

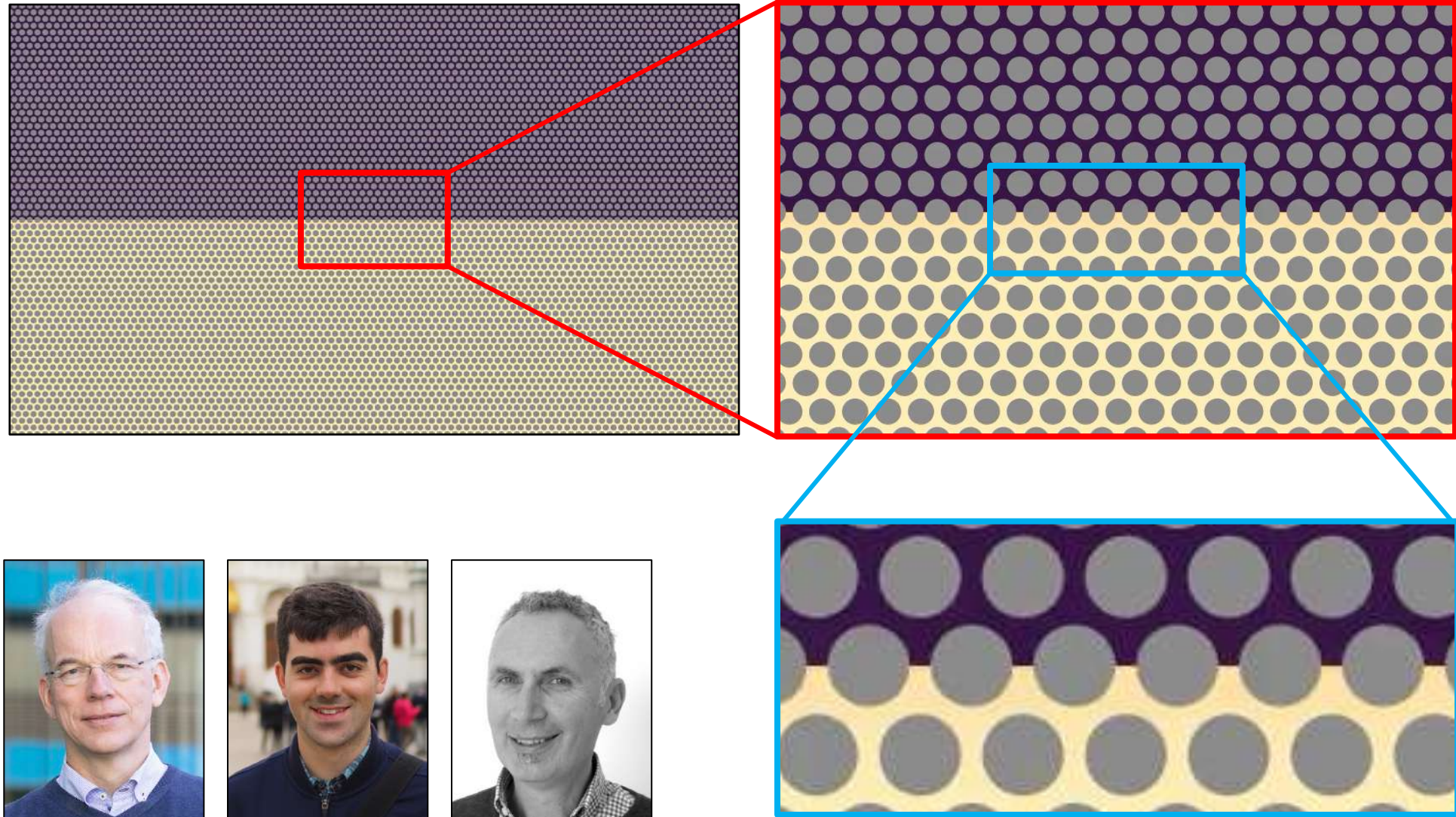








[arxiv.org/abs/2310.04068](https://arxiv.org/abs/2310.04068)



T32:1: Rayleigh-Taylor instability in confined porous media, **De Paoli**, Howland, Verzicco & Lohse

Mon, Nov. 20<sup>th</sup>  
16:25, Room 158AB





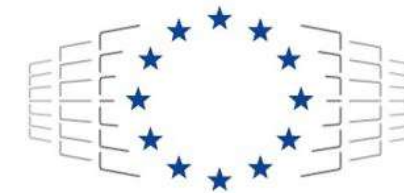
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**PREPRINT**



**EuroHPC**  
Joint Undertaking

This project has received funding from the European Union's Horizon Europe research and innovation programme under the Marie Skłodowska-Curie grant agreement MEDIA No. 101062123.



**Funded by  
the European Union**



High-resolution images, movies and slides are available upon request to [m.depaoli@utwente.nl](mailto:m.depaoli@utwente.nl)